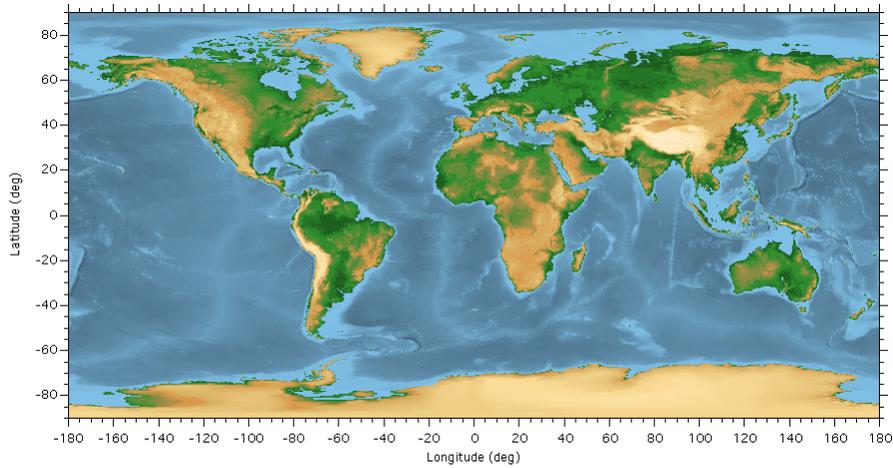


GG 611 Big Gulp
Fall 2014

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POST 804
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Gravity, the Geoid, and Mantle Dynamics

Lecture: Plate Tectonics & Mantle Convection



Francis Bacon (1561-1626)



Ben Franklin (1706-1790)



The Known World, ~1700

**Was the Atlantic once
closed?**

If so, how could this be?

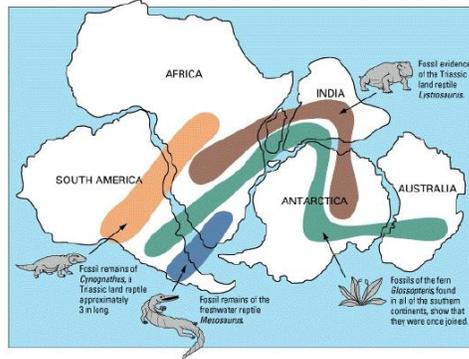


Alfred Wegener
(1880-1930)

The two sides of the Atlantic:

- Matching Shapes
- Matching Rock Types
- Matching Fossils
- Patterns of Uplift

Fossil Distributions



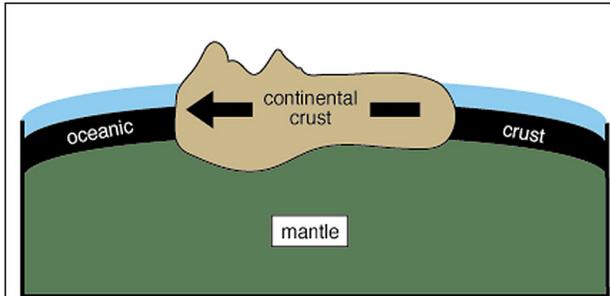
Wegener's Idea: *The Origin of Continents and Oceans* (1915)

Continental Drift:

The continents plowed through the oceanic crust

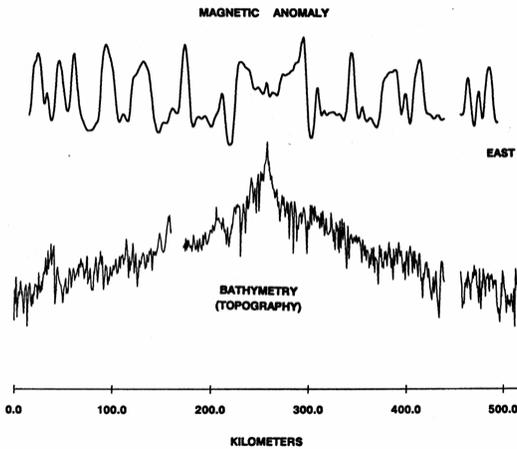
Problem:

Wegener did not describe the forces that caused continents to move!



Wegener's proposal that continents plowed through oceanic crust was not accepted by other geologists.

1950s - Magnetic Surveys of the Seafloor



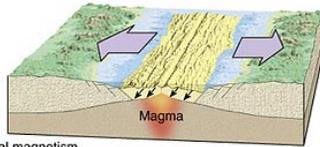
Magnetic Anomaly:
symmetric about ridge

Seafloor Bathymetry:
descends away from ridge

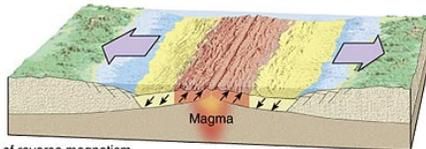
1960s: The development of **Plate Tectonics**

Harry Hess, 1962: "A History of the Ocean Basins"

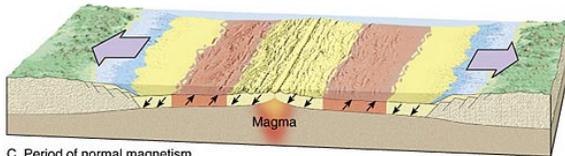
Seafloor spreading allows plates to move



A. Period of normal magnetism



B. Period of reverse magnetism

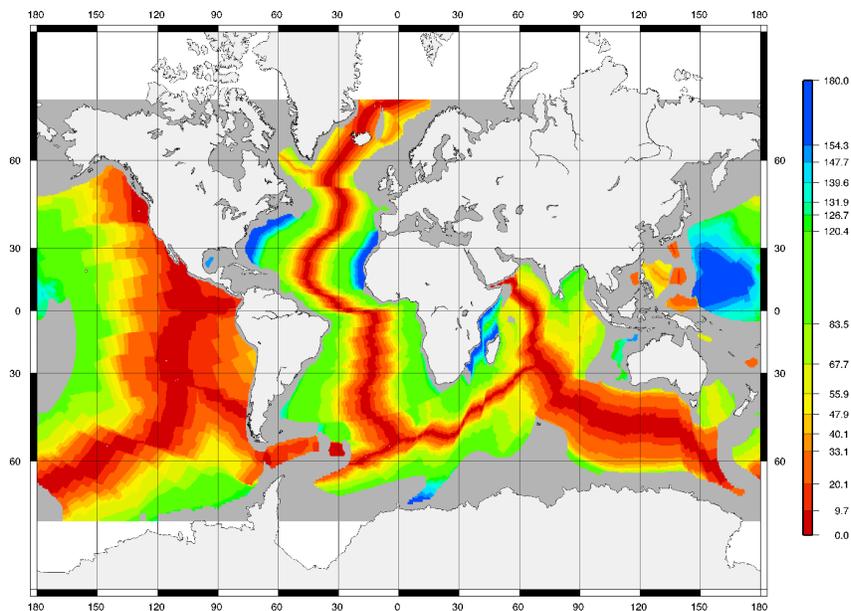


C. Period of normal magnetism

New seafloor is being created as the seafloor spreads

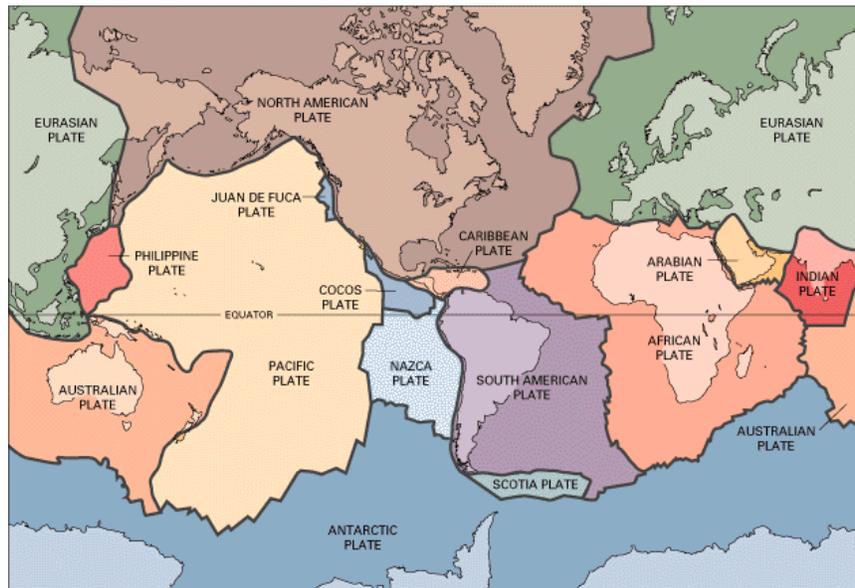
The magnetic field is "frozen" in the newly-created seafloor.

Mapping Seafloor Ages



Oldest Seafloor = 180 Million Years!

Earth's Tectonic Plates



Biggest Plate: Pacific Plate (10,000 km across)
 Smallest Plate: Juan de Fuca (300 km across)
 All sizes in between!

How Fast do the Plates Move?

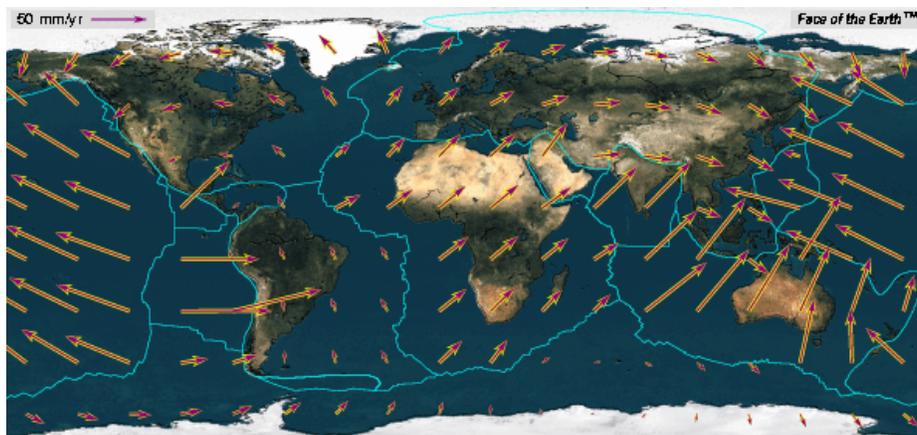
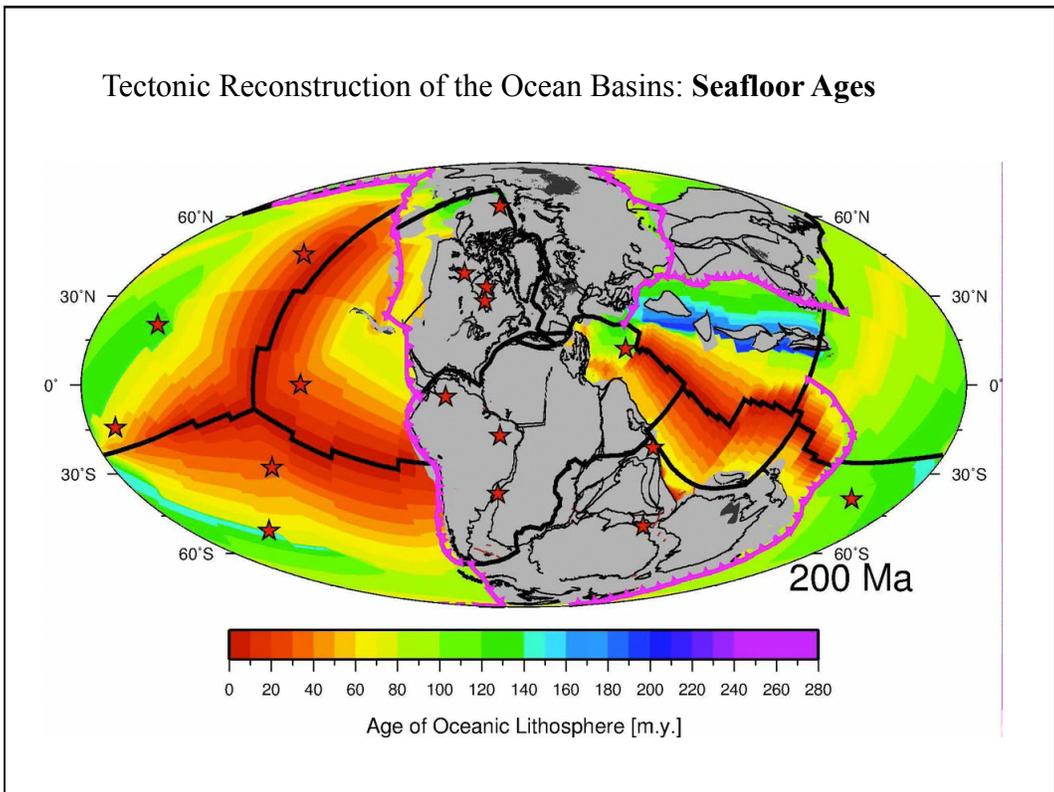
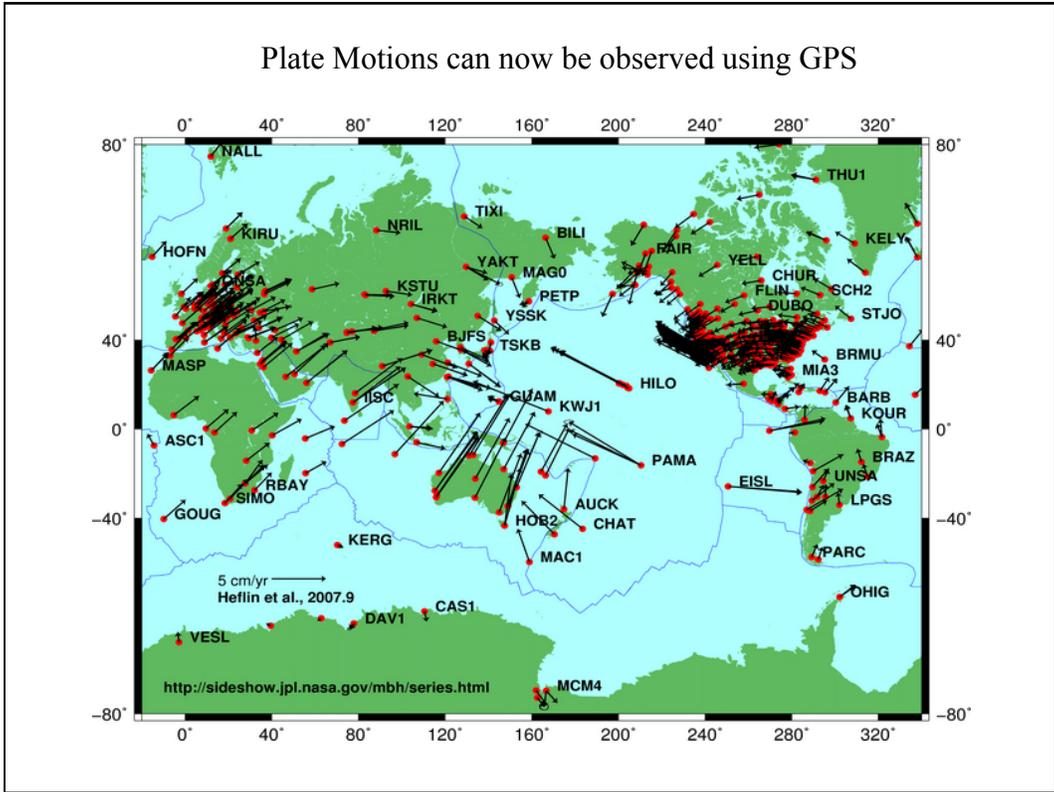
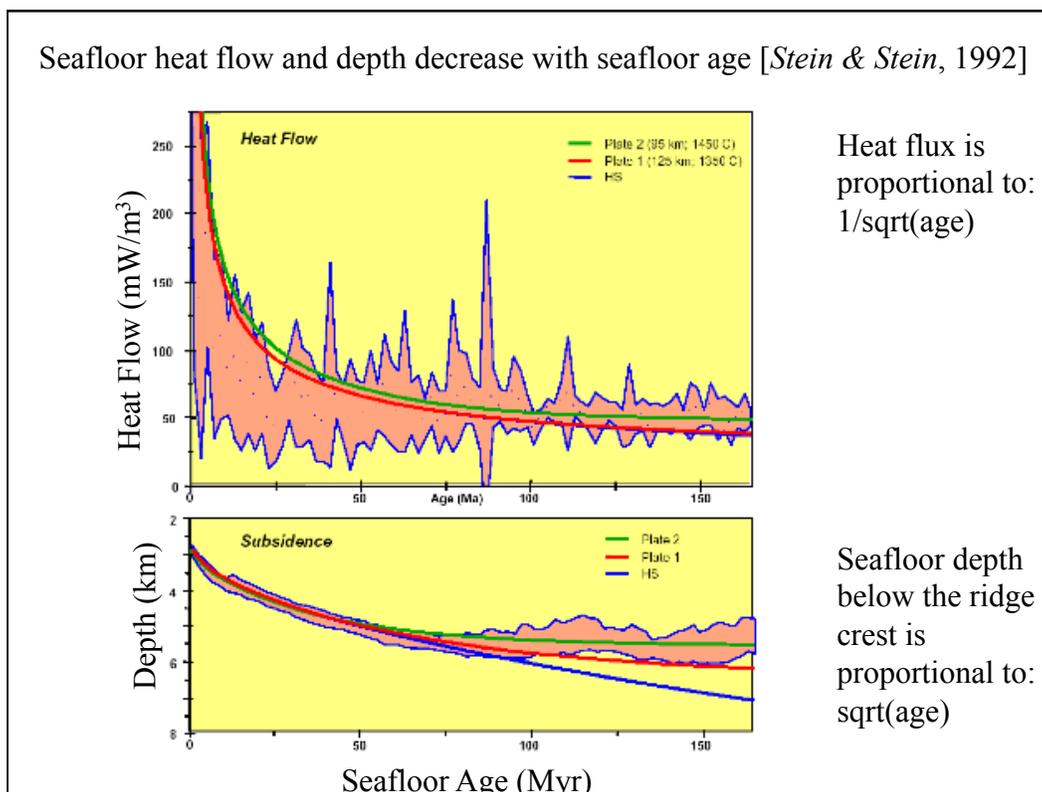
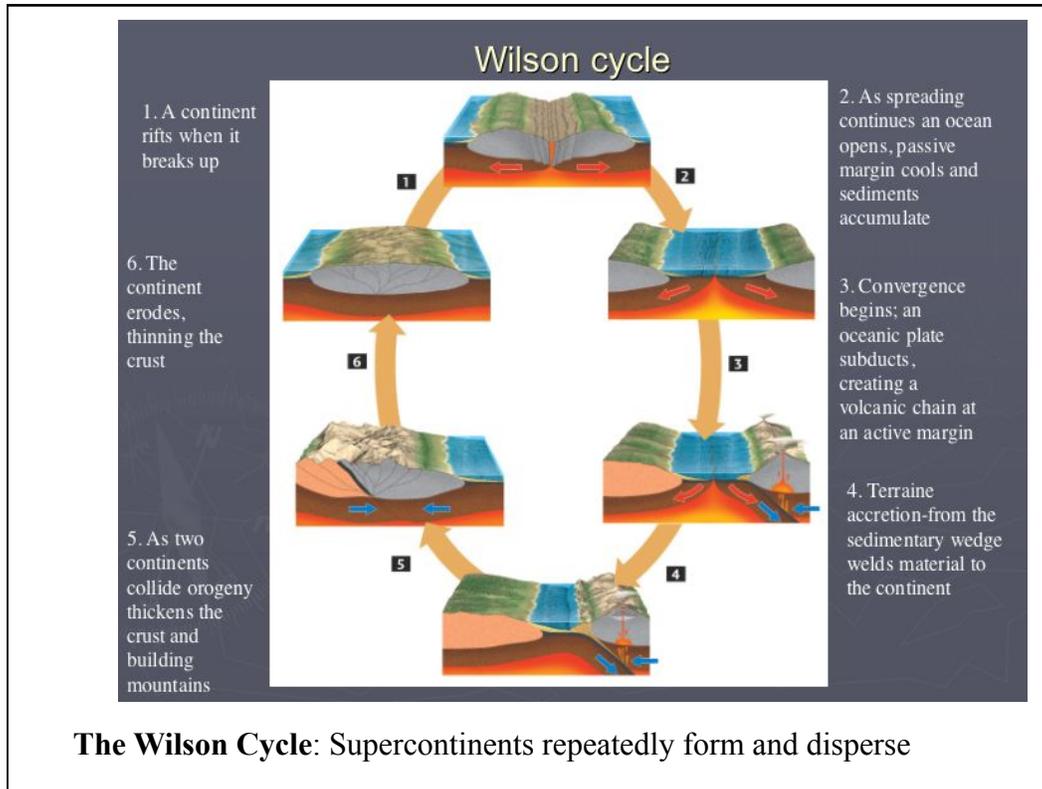
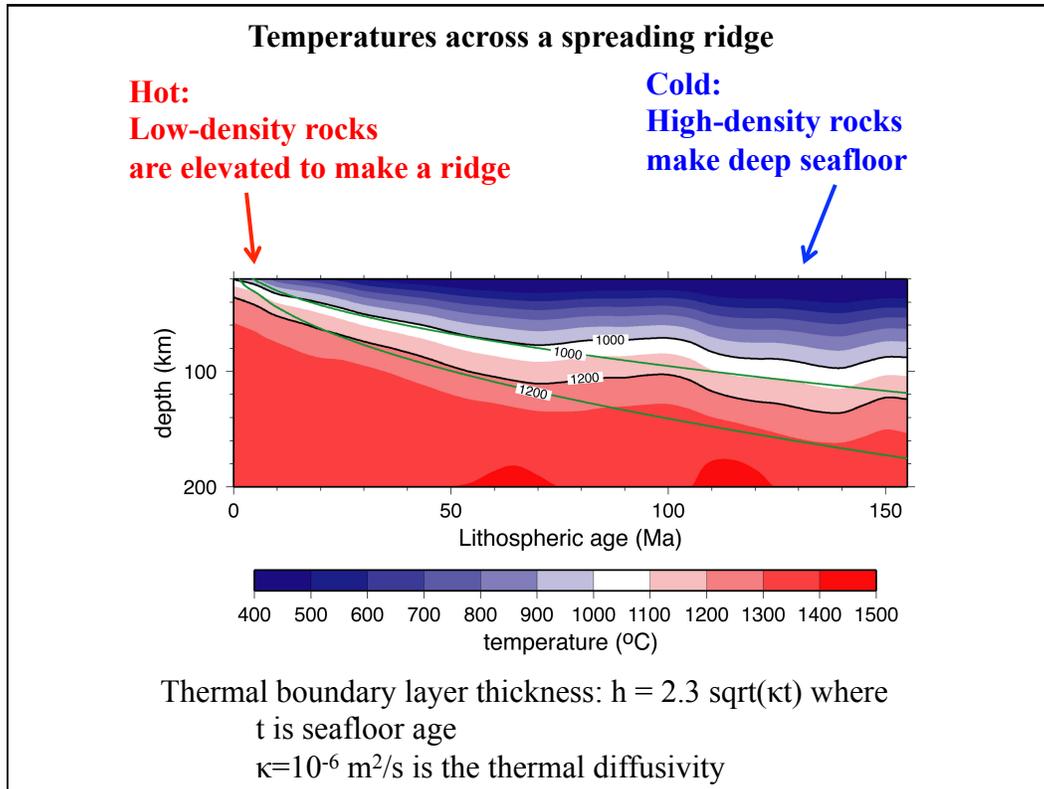


Plate Speeds

Atlantic Basin:	1-2 cm/yr
Indian Basin:	3-7 cm/yr
Pacific Basin:	5-10 cm/yr







**What drives the plate motions?
WHY do the plates move?**

The Earth loses heat through CONVECTION

Examples of Convection

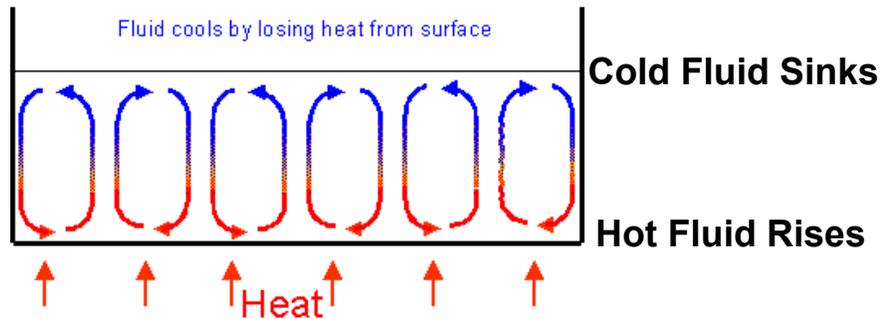


Lava lamp



Boiling water

How Convection Works

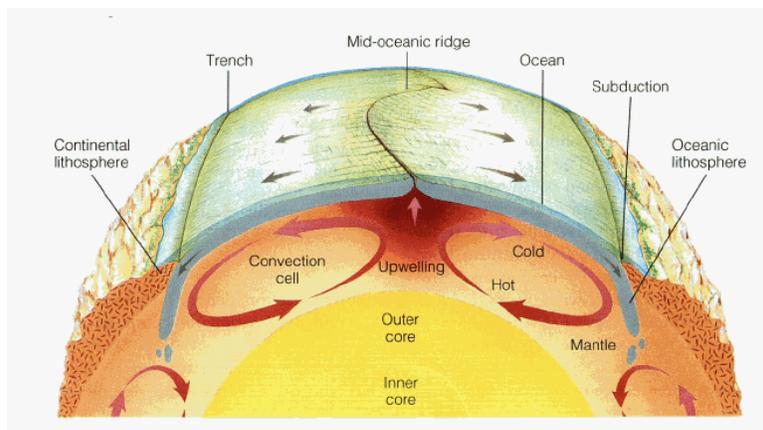


Convection cell

Warm, low density fluid rises
Cool, high density fluid sinks

**Convection is the
fastest way for the
Earth to get rid of heat**

Mantle Convection and Plate Tectonics



UPWELLING beneath spreading ridges
DOWNWELLING beneath subduction zones
THE PLATES surface expression of
 mantle convection

The Rayleigh Number is a dimensionless parameter that measures the vigor of convection:

$$Ra = \frac{\rho g \alpha \Delta T D^3}{\kappa \eta}$$

ρ = density (3300 kg/m³)

g = gravity (10 m/s²)

α = thermal expansivity (3×10^{-5} K⁻¹)

ΔT = Temperature contrast across mantle (3000 K)

D = Depth of Mantle (2860 km)

κ = Thermal diffusivity (10^{-6} m²/s)

η = Mantle viscosity (10^{21} Pa s)

Using these parameters: $Ra \sim 7 \times 10^7$

Convection occurs if Ra is larger than a critical value Ra_{cr}

For convection in a layer, $Ra_{cr} = 657$

Mantle Convection: How Vigorous?

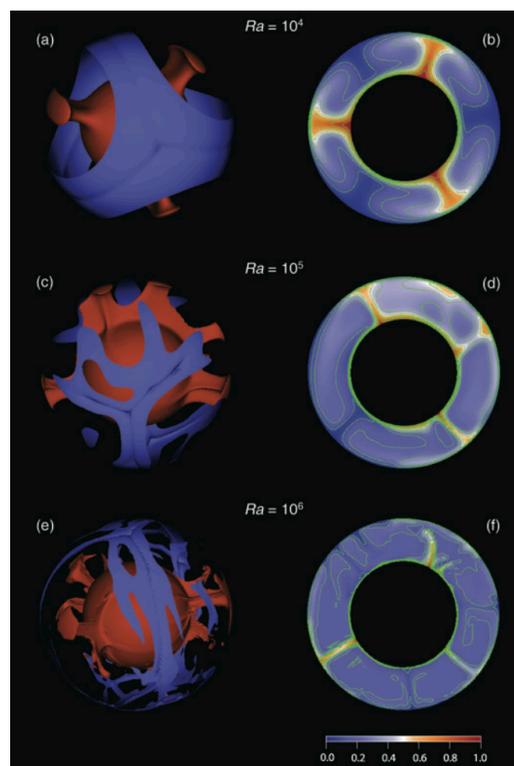
Deschamps et al., 2010

Style and vigor of
convection changes
with Ra

Boundary layer
thickness $h \sim Ra^{-\frac{1}{3}}$

Plate velocity $v_p \sim Ra^{\frac{2}{3}}$

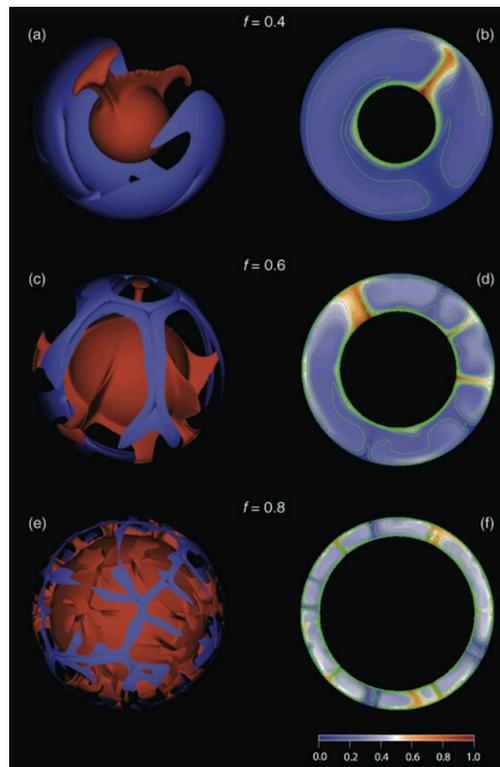
Mantle heat flow $Q \sim Ra^{\frac{1}{3}}$



Mantle Convection: Effect of Core Size $Ra = 10^5$

Deschamps et al., 2010

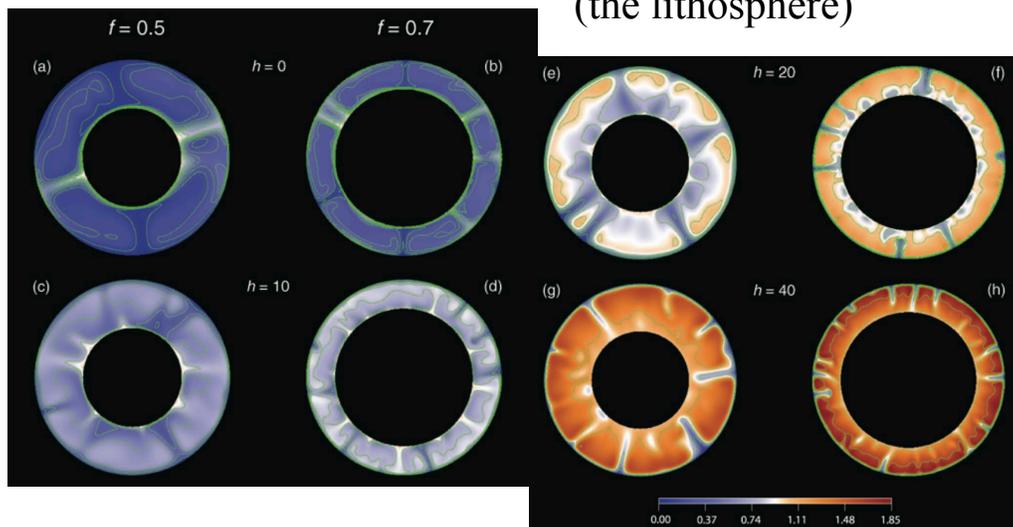
Convection cells form
with an aspect ratio
greater than 1

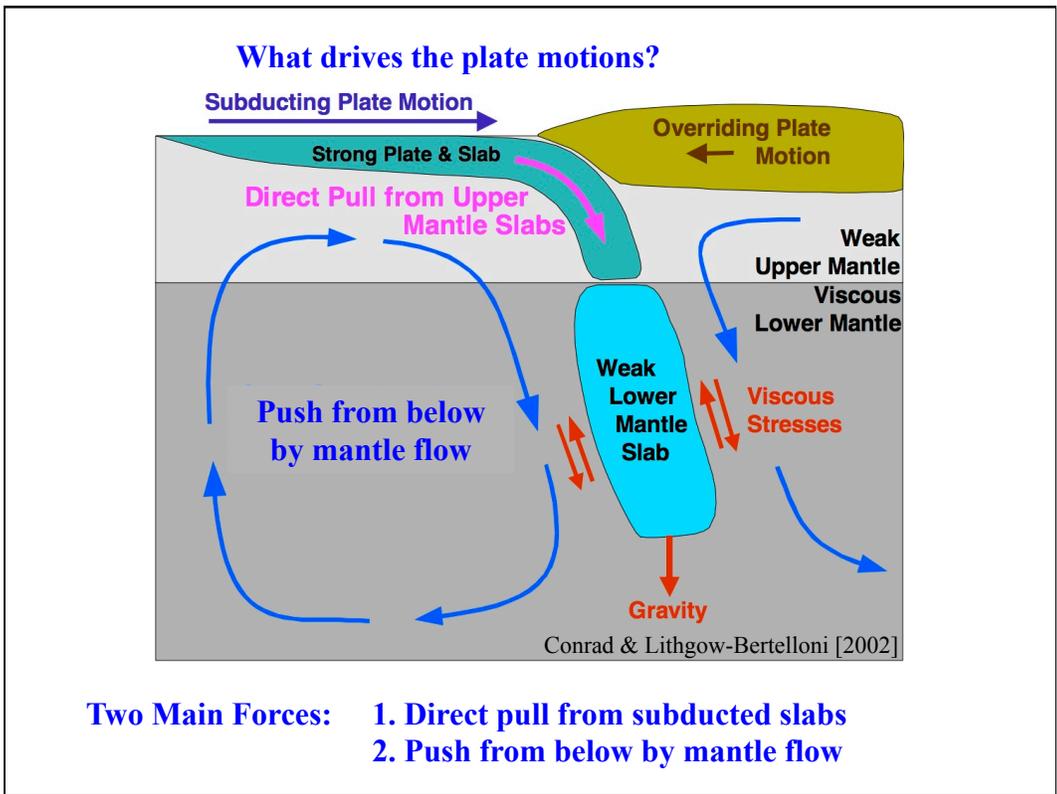
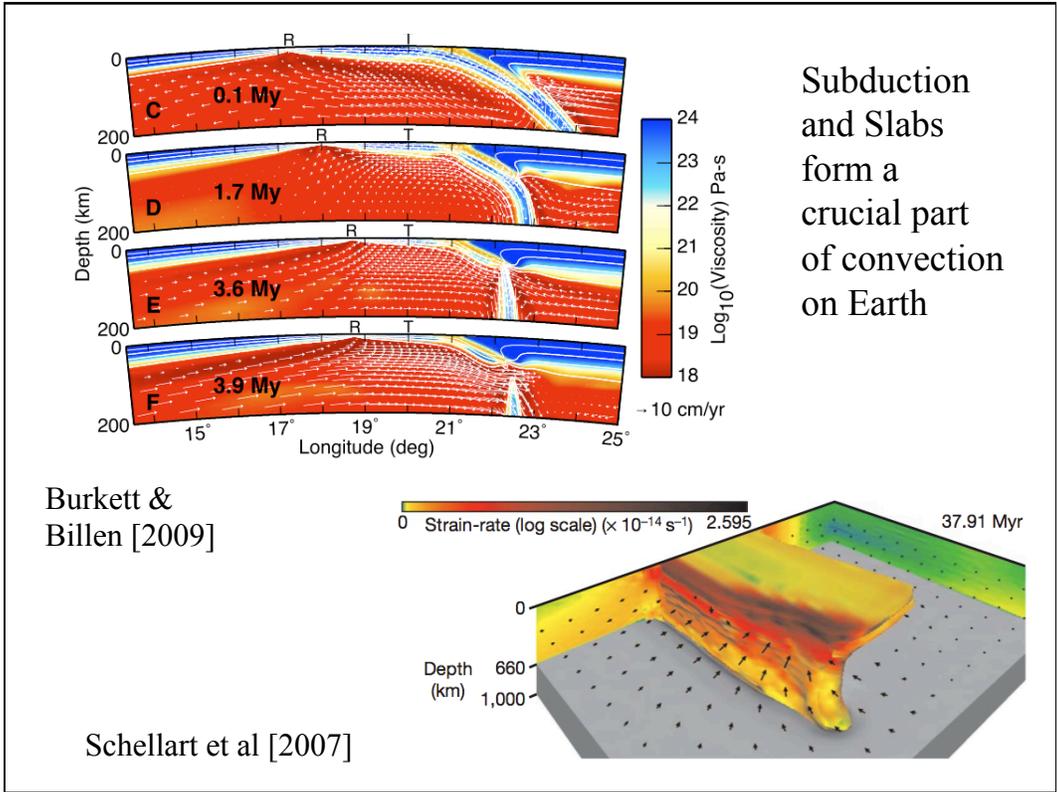


Mantle Convection: Effect of Internal Heating

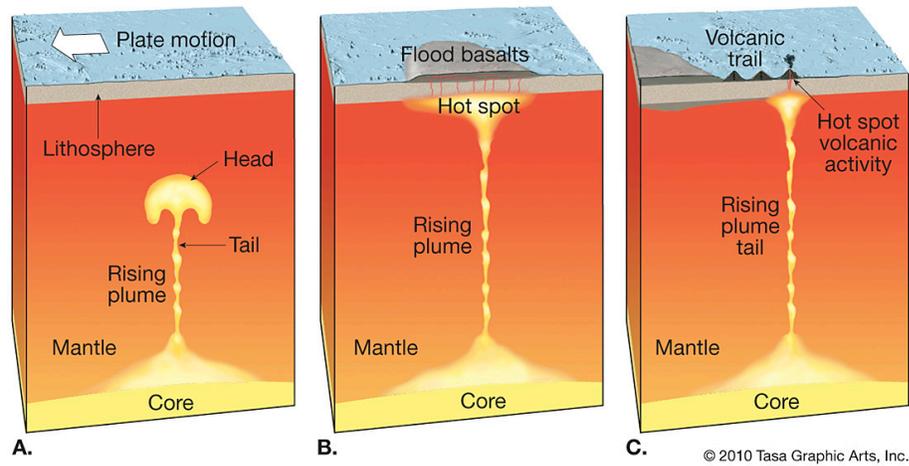
Deschamps et al., 2010

Internal heating increases
the importance of the
top boundary layer
(the lithosphere)

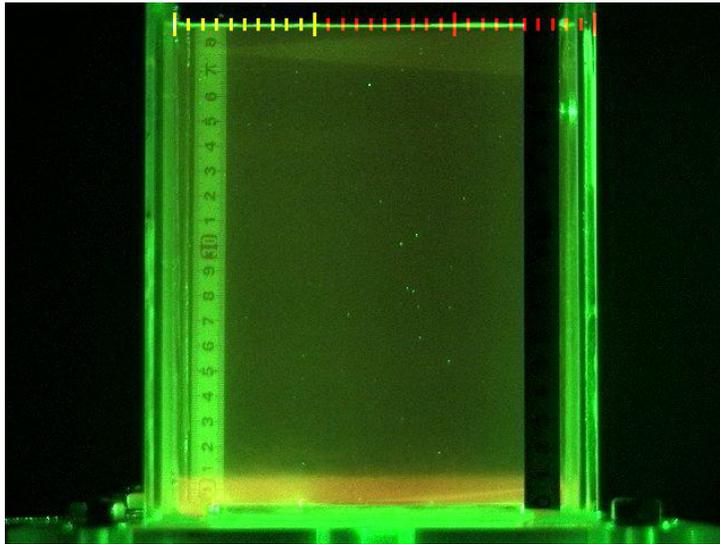




Mantle Plumes and Hotspots

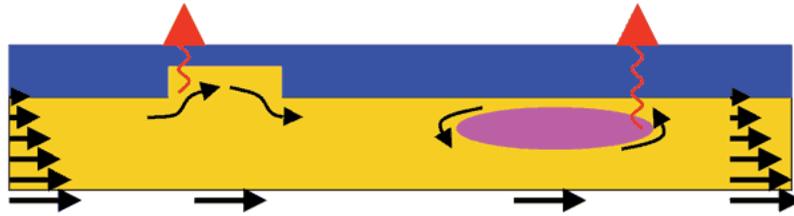


Mantle plumes explain volcano chains within plate interiors.
Plumes rise from the lower boundary layer of convection.



A Plume Experiment in Corn Syrup

Shear-Driven Upwelling



Like a plume moving sideways:
Shearing beneath the tectonic plates produces
upwelling and volcanism (if viscosity is heterogeneous)

Explains some minor volcanism [*Conrad et al.*, 2011].

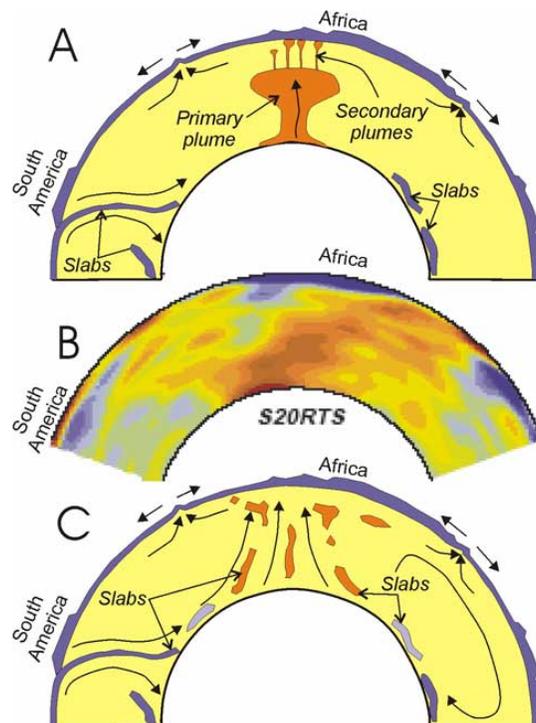
Superplumes?

Two giant slow
seismic anomalies
are observed
beneath Africa &
the Pacific

Tomography across Africa

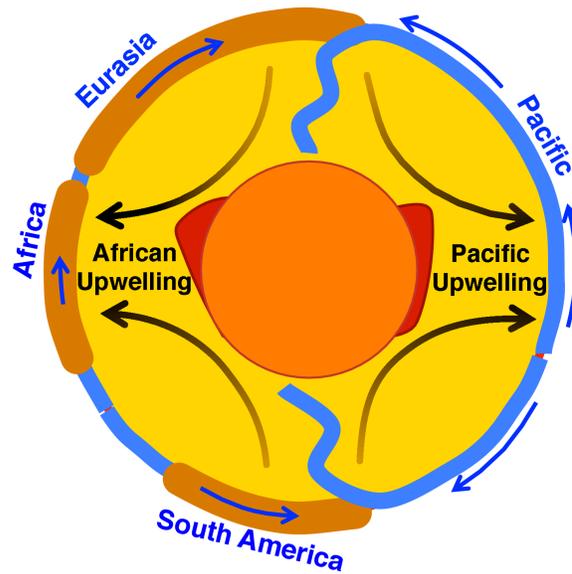
These may
represent major
upwellings in the
mantle.

Plume Cluster?



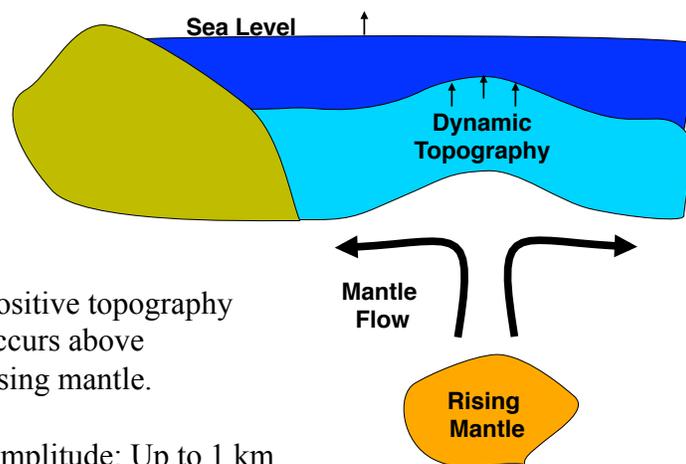
The superplumes appear to be stable.

They may organize the pattern of flow in the mantle:



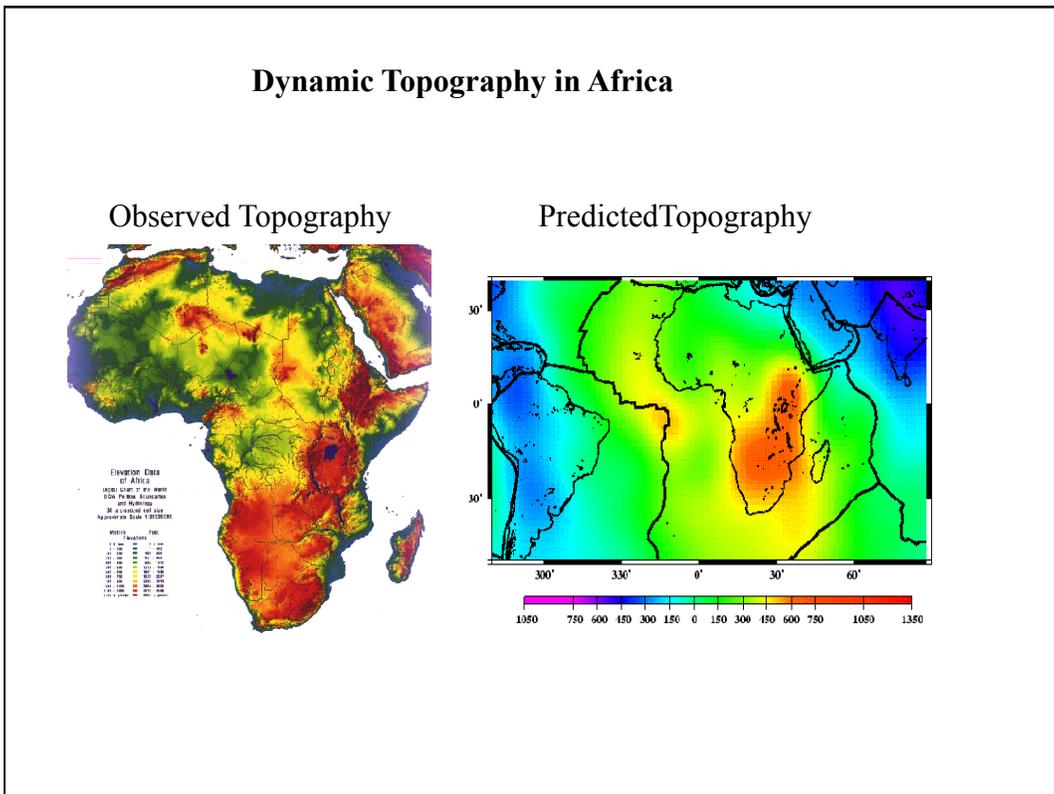
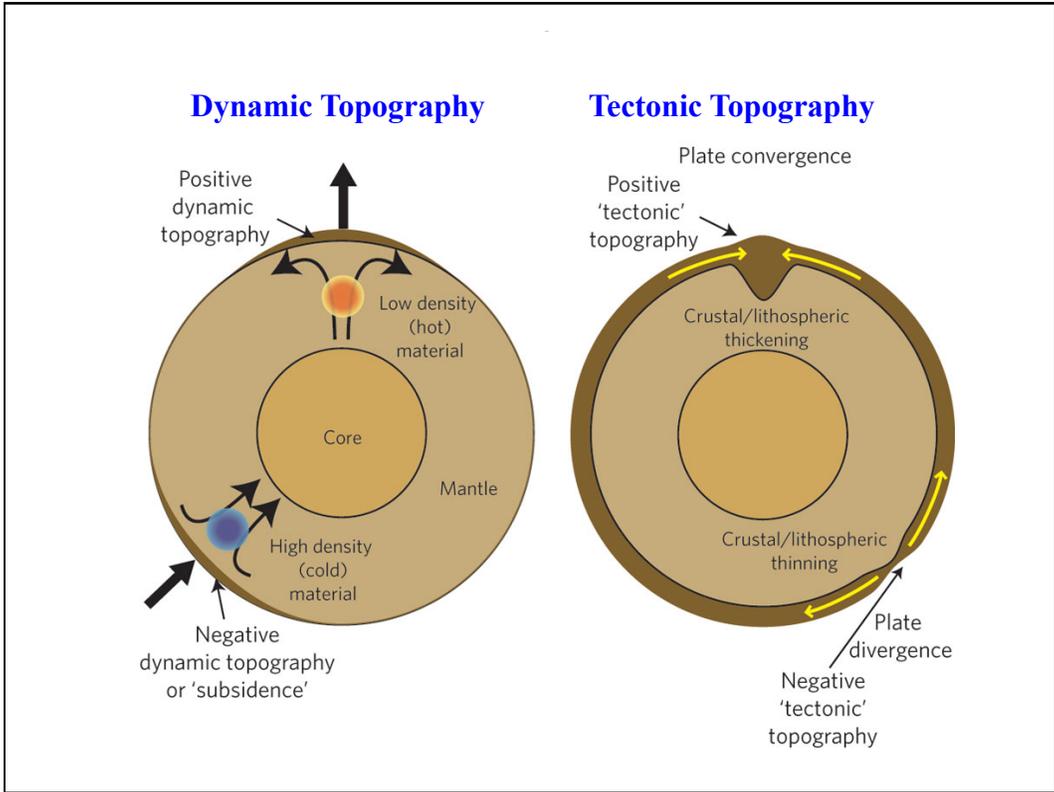
Conrad et al. [2013]

**Mantle Flow causes surface uplift and subsidence.
This topography is called “Dynamic Topography”.**



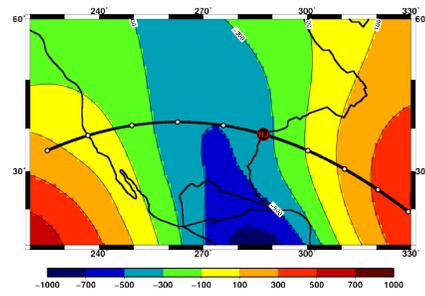
Positive topography occurs above rising mantle.

Amplitude: Up to 1 km



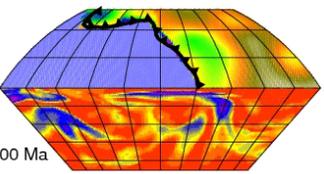
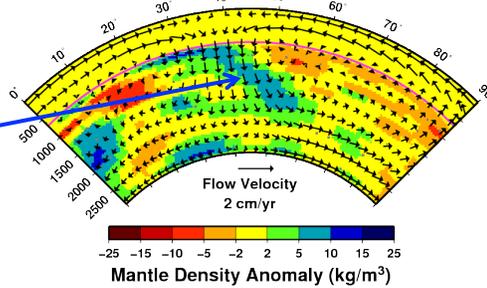
Dynamic Topography in North America

Downwelling beneath the Eastern US depresses the surface

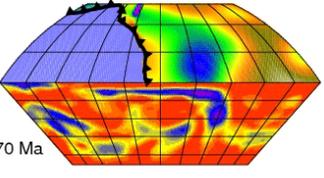


Dynamic Topography at 0 Ma (km)

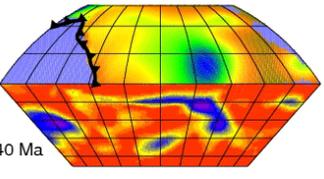
Farallon Slab



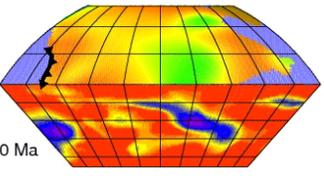
100 Ma



70 Ma



40 Ma



0 Ma

Dynamic topography caused the Western Interior Seaway



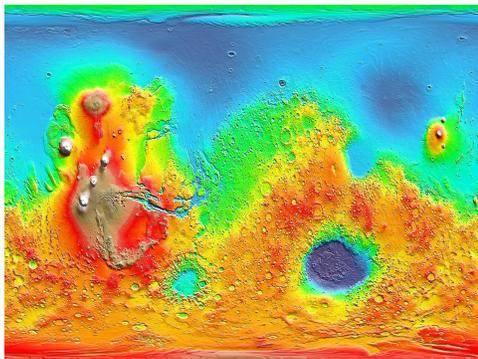
Maximum extent of the Western Interior Seaway about 100 million years ago

Plate Tectonics and Convection on other Planets?

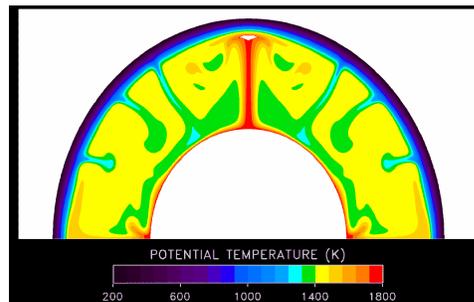
Moon, Mars, Venus, Mercury: Surfaces are much older than Earth's: Probably no plate tectonics

Instead, mantle convection may take a different form

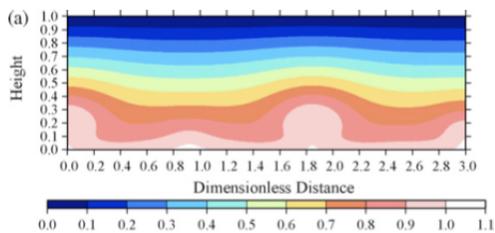
Mars Topography



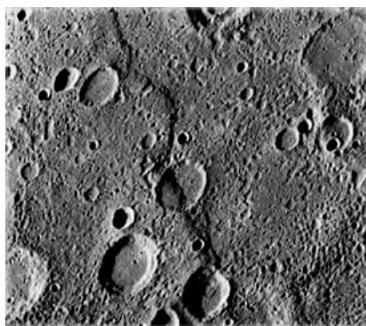
Model of Mars Convection



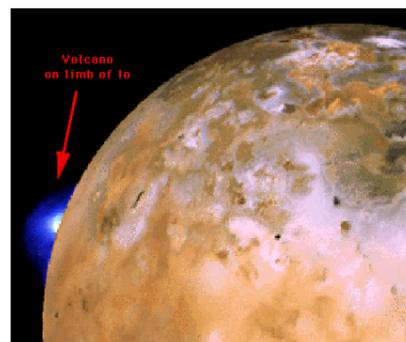
Mercury: Low Ra

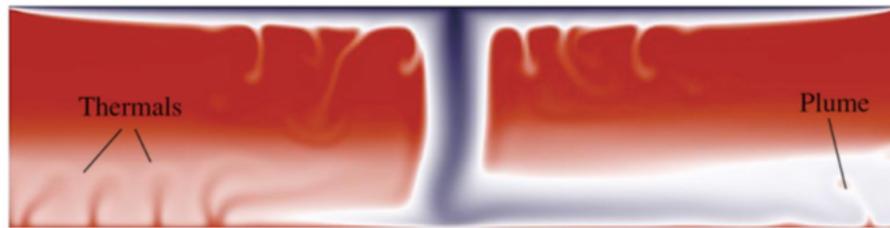


Redmond & King 2007



Io: High Ra

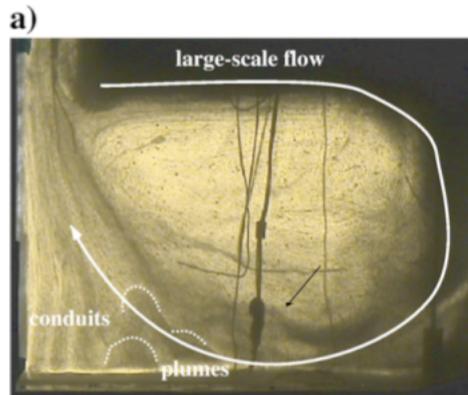




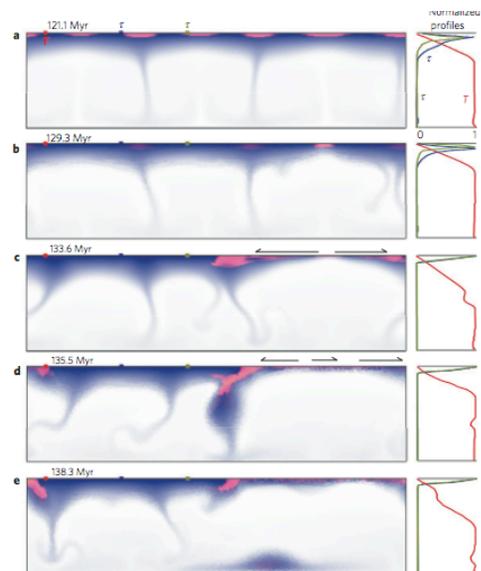
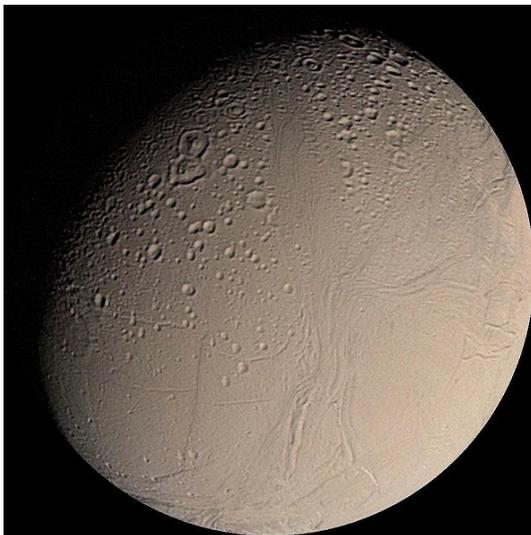
Venus:

No plate tectonics,
but the entire lithosphere
sometimes sinks into
the mantle, resurfacing
the entire planet.

Robin et al., 2007

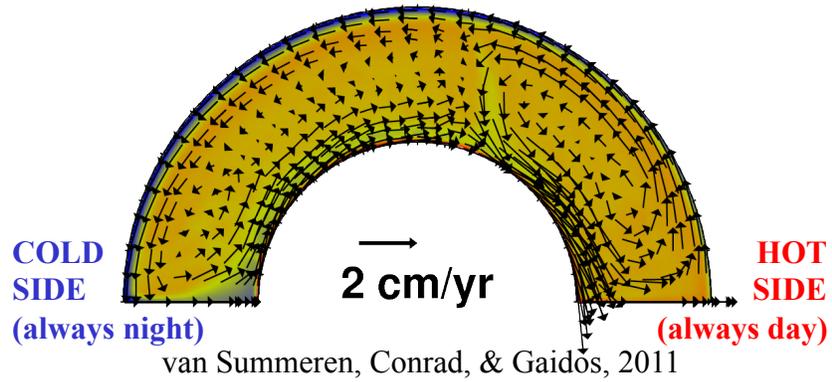


Enceladus: Convection in solid ice



Exoplanets: Many different styles!

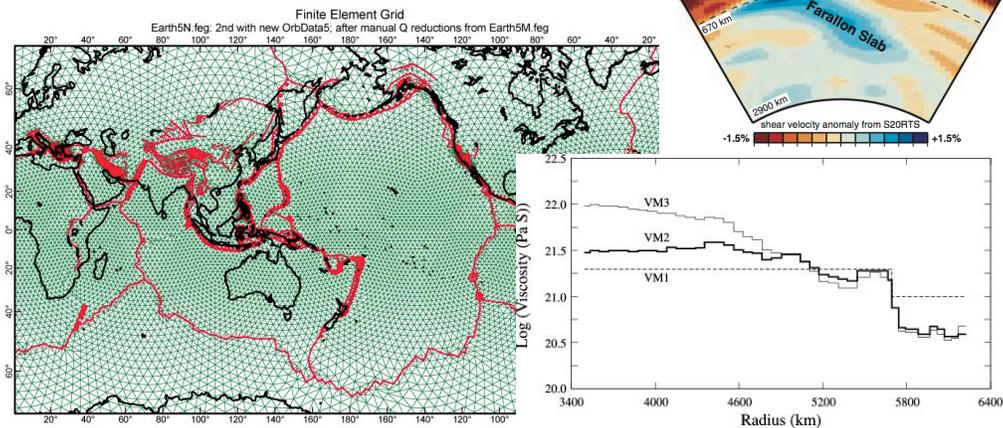
Tidally-locked example



Numerical Models of Mantle Convection

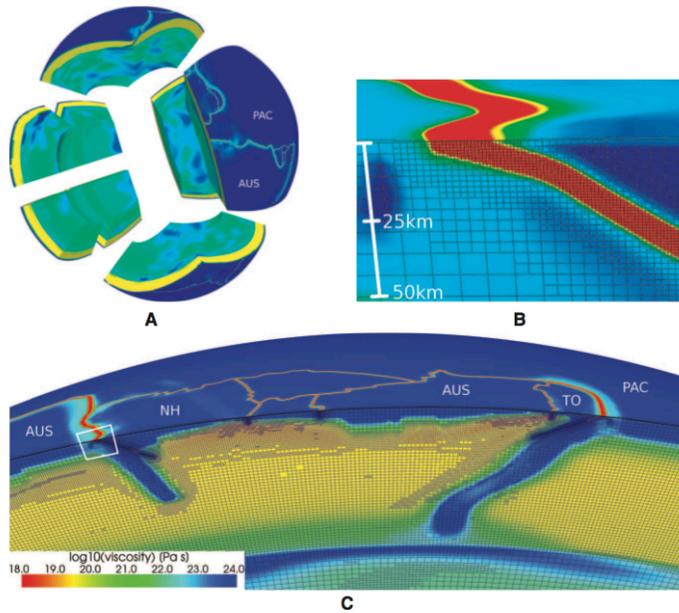
Input:

1. Mantle Densities
(from seismic tomography)
2. Mantle Viscosity Structure
(from postglacial rebound)
3. A Finite Element Code

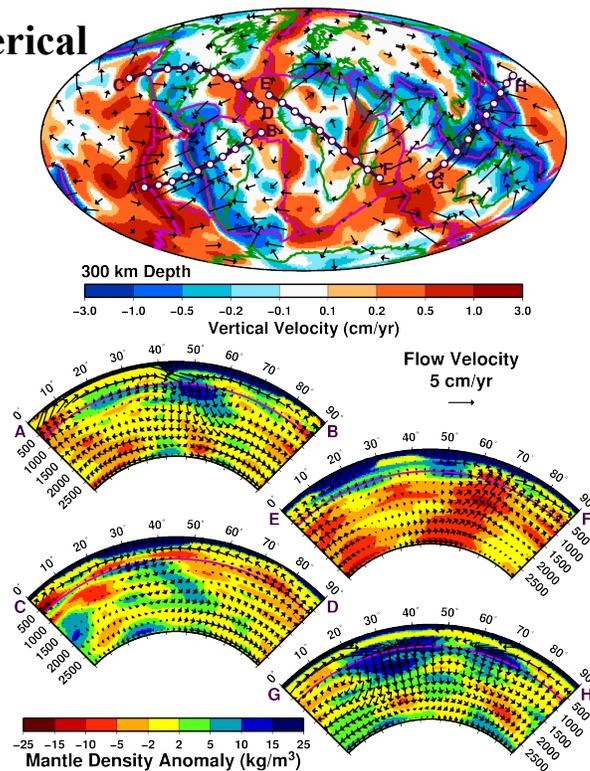


State of the Art: Adaptive Mesh Refinement & Slabs

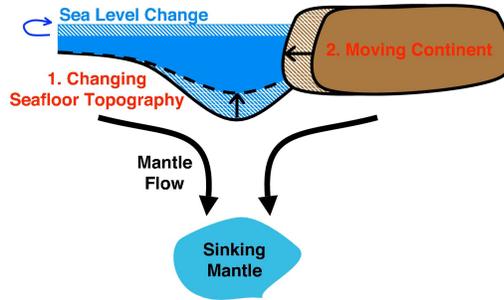
Stadler et al.,
Science 2010



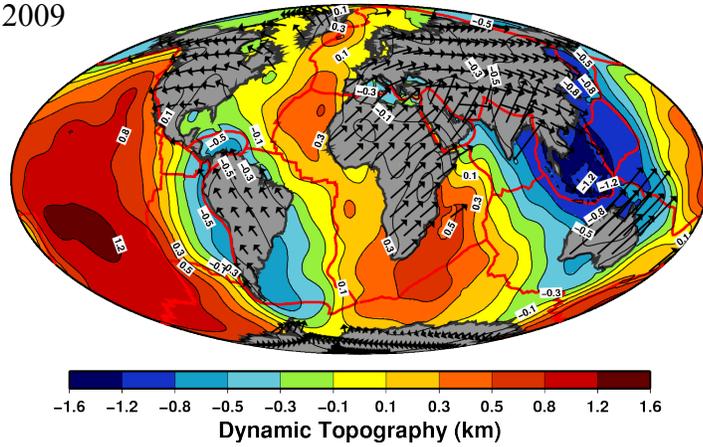
Predictions of Numerical Convection Models: Mantle Flow Field



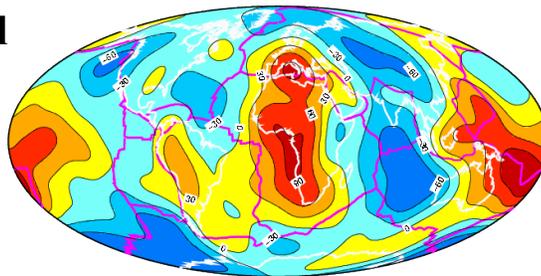
Predictions of Numerical Convection Models: Dynamic Uplift



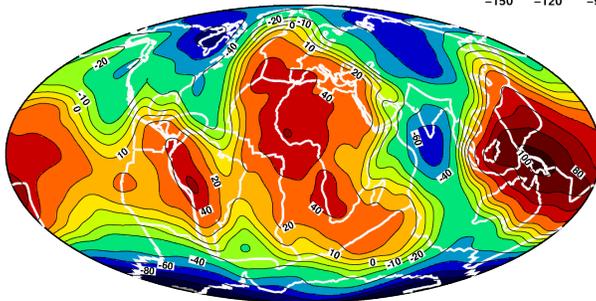
Conrad & Husson, 2009



Predictions of Numerical Convection Models: Earth's Shape (the Geoid)

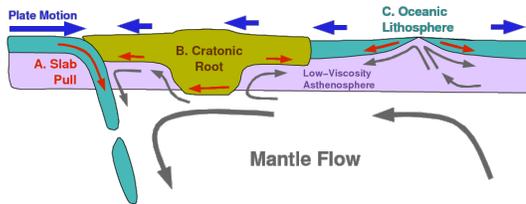


Observed Geoid (grimg)

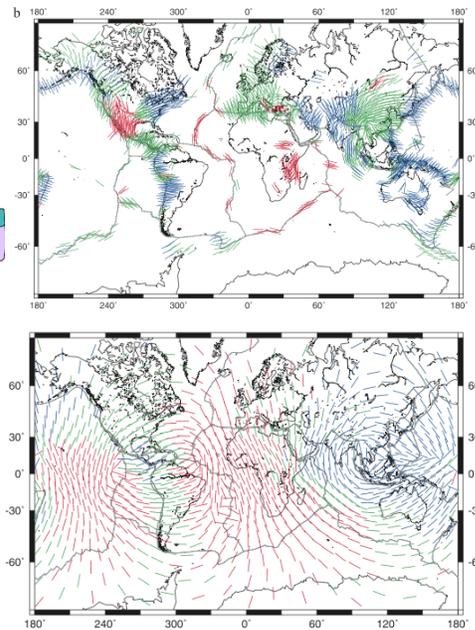


Geoid Height (m)

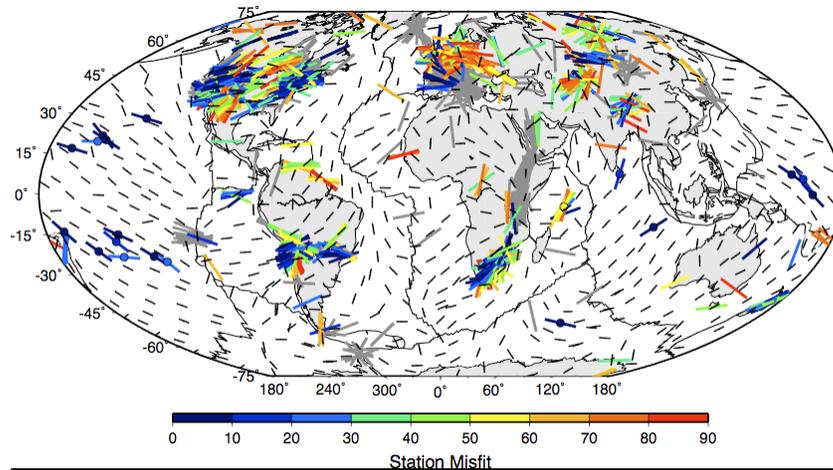
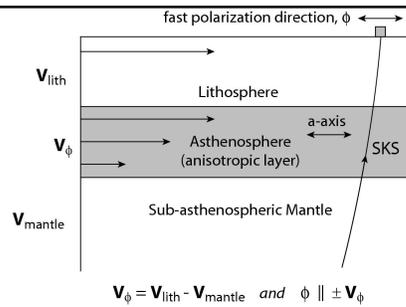
Predictions of Numerical Convection Models: Lithospheric Stresses



Lithgow-Bertelloni & Gynn, 2004

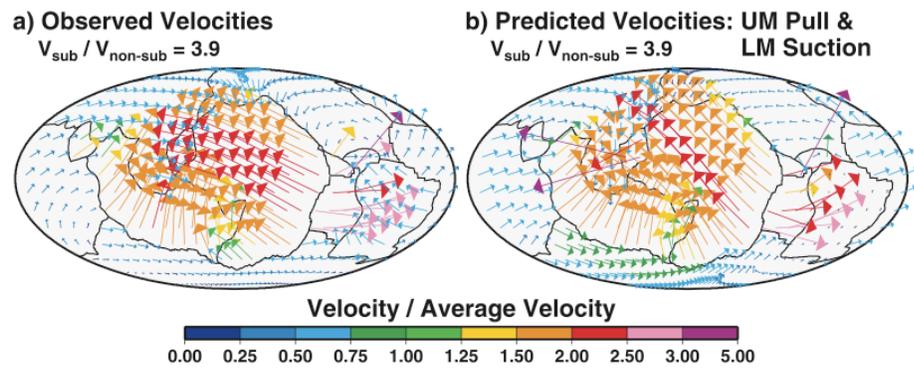


Predictions of Numerical Convection Models: Anisotropic Fabric



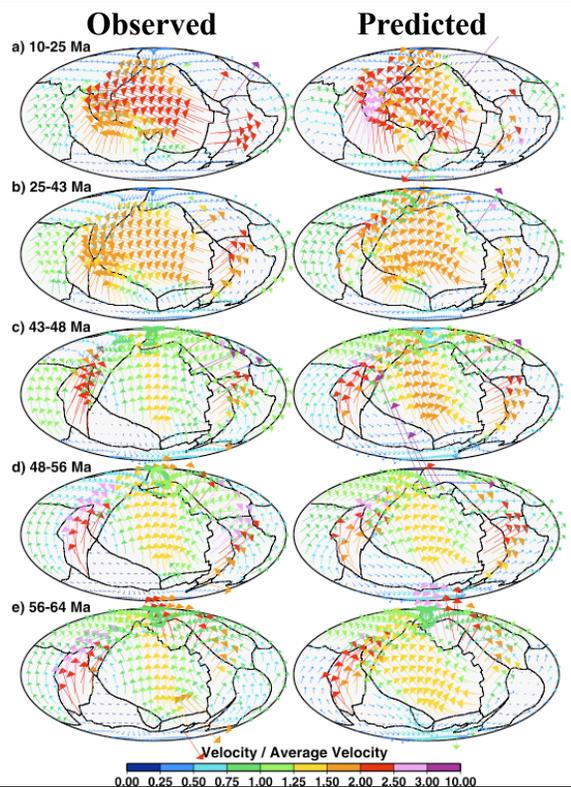
Conrad et al., 2007

Predictions of Numerical Convection Models: Plate Motions



Conrad & Lithgow-Bertelloni [2004]

Predictions of Numerical Convection Models: Time-Dependence of Plate Motions



Conrad & Lithgow-Bertelloni [2004]

Models of Earth's Thermal History: Heat flow $\sim Ra^{1/3}$

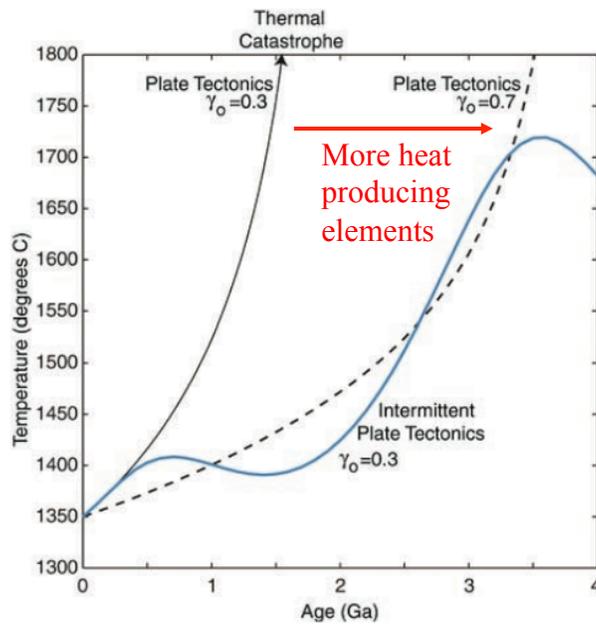
We know the present cooling rate from Earth's heat flow

We estimate the volume of heat-producing elements

Problem: models give a hot start to the earth at 1.5 Ga (!)

Solutions:

- Deep reservoirs of heat producing elements
- Plate tectonics is somehow less efficient



Silver & Behn [2008]